

https://doi.org/10.1093/pnasnexus/pgac220 Advance access publication date: 14 October 2022 Research Report

The lifetime of charged dust in the atmosphere

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Edited By: Doraiswami Ramkrishna

Abstract

Wind-blown dust plays a critical role in numerous geophysical and biological systems, yet current models fail to explain the transport of coarse-mode particles (>5 μ m) to great distances from their sources. For particles larger than a few microns, electrostatic effects have been invoked to account for longer-than-predicted atmospheric residence times. Although much effort has focused on elucidating the charging processes, comparatively little effort has been expended understanding the stability of charge on particles once electrified. Overall, electrostatic-driven transport requires that charge remain present on particles for days to weeks. Here, we present a set of experiments designed to explore the longevity of electrostatic charge on levitated airborne particles after a single charging event. Using an acoustic levitator, we measured the charge on particles of different material compositions suspended in atmospheric conditions for long periods of time. In dry environments, the total charge on particle material and charge polarity. However, exposure to UV radiation could both increase and decrease the decay time depending on polarity. Our work suggests that the rate of charge decay on airborne particles is solely determined by ion capture from the air. Furthermore, using a one-dimensional sedimentation model, we predict that atmospheric dust of order 10 μ m will experience the largest change in residence time due to electrostatic forces.

Keywords: charged dust, acoustic levitation, atmospheric electricity

Significance Statement:

Wind-blown dust provides nutrients to the oceans and has significant impacts on the Earth's radiative balance. Long-range transport should involve only very small particles. Yet, observations repeatedly find large dust grains thousands of kilometers away their sources. Because dust can charge during lofting, electrostatic forces may explain long residence times of coarse-mode dust. However, the temporal stability of electrostatic charge on dust is unknown. We bridge this knowledge gap by monitoring the charge decay on acoustically levitated electrified dust particles in various environments. We show that charge decay is primarily driven by ion capture. Particles remain charged for days or weeks, suggesting that electrostatic forces are important for the transport of dust with diameters of order $10 \,\mu$ m.

Introduction

Atmospheric dust is an important component of local and global climate systems. Mineral dust is the most abundant aerosol type by mass and is emitted into the atmosphere at a rate of up to 5,000 Tg/y (1, 2). Small, localized events, like volcanic eruptions or wild fires also contribute to Earth's dust budget [volcanoes, for instance, inject ~13 Tg of ash into the atmosphere every year (3)]. At regional scales, dust emissions from wildfires, volcanoes, and dust storms may rapidly and dramatically alter local atmospheric dust budgets. These acute increases in solid mass loading represent hazards to populations, natural environments, and infrastructure (4, 5). On a planetary-wide scale, atmospheric dust interacts with short- and long-wave radiation (6), tuning the Earth's energy balance (7). In turn, this modulation has profound impacts on sea and land surface temperatures, atmospheric circulation, and weather (8, 9). Solid particles may also serve as cloud condensation and ice nuclei, influencing the formation of clouds and impacting precipitation rates (10). Airborne particles have the ability to carry nutrients, toxins, and bacteria across long distances. Indeed, Saharan desert dust has been recognized as an important source of phosphorous for the Amazon rain forest (11). Similarly, lofted dust can modify atmospheric chemistry by providing reactive substrates for various chemical species (12). Beyond Earth, atmospheric dust likely influences surface processes on Venus, Mars, Io, Titan, and Gliese J1214b (13–16).

Since the 1970s, evidence indicates that climate models fail to accurately represent the quantity of coarse-mode particles (with diameters $D > 5\,\mu$ m) in the atmosphere (17, 18) by a factor of 4 (18). Although dust transport models predict that coarse dust settles rapidly, observations consistently find large particles at distances well beyond the maximum ranges estimated numerically. For instance, Denjean et al. showed that the effective diameter of coarse-mode Saharan dust over the Mediterranean remained unchanged for up to a week after being lofted into suspension

Competing Interest : The authors declare no competing interest. Received: June 21, 2022. Accepted: October 3, 2022

OXFORD UNIVERSITY PRESS

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(19). Likewise, modeling shows that particles in the range of 20 to 30 μ m should settle out of the Saharan Air Layer in approximately 1.5 to 3 days. Yet, grains with these diameters have been detected over the Caribbean after 4,000 km and 5 days of transport (17). Maring et al. (20) found that the distribution of dust over the Canary Islands requires that particles experience an upward velocity of ~0.33 cm/s imparted by some unknown process.

Recently, electrostatic forces have been invoked to account for extended residence times of dust in the atmosphere (21). Particles may charge as they are injected into the atmosphere through aeolian action, splashing, chemical processes, fracture, and comminution (22, 23). Charging mechanisms include fracto- and triboelectric charging (24, 25), radioactive decay (26), and gas ionization (27). Surprisingly, a number of investigations have demonstrated that airborne particles may remain charged even at great distances from where they emanate. For example, high space charge densities were found in an ash cloud 1,200 km from its source at the Eyjafjallajökull volcano (28), and Saharan dust over Scotland can carry an edge charge density several times larger than that of typical stratiform clouds (29).

Accurately assessing the degree to which electrostatic forces influence the transport of large particles requires better constraints on (i) the magnitude of the electric fields within atmospheric dust layers, (ii) the mechanisms by which particles charge before and after being lofted, and (iii) the ability of electrified grains to retain charge once airborne. Improvements to our understanding of the meso- to macroscale electrostatic characteristics within dusty environments may come from a combination of airborne and ground-based measurements, complemented by numerical modeling (18, 28, 30, 31). However, elucidating charge evolution on particles fundamentally requires long-term measurements at the grain scale. Such measurements are difficult to do in-situ and, until recently, laboratory experiments have been unable to isolate charged grains from other surfaces to mimic airborne transport.

Here, we investigate the longevity of charge on isolated particles of various compositions suspended in the air. We implement a non-contact charge measurement technique that can monitor charge decay of a single acoustically levitated particle for days to weeks. Unlike previous investigations, which allowed charge to leak away across contact points, the present experiments characterize charge loss occurring only at the particle-gas interface. Our results suggest that lofted particles can retain charge for weeks, with the decay rate depending only on environmental factors like relative humidity, charge polarity, and irradiation. Relative humidity (RH) decreases the half-life $(t_{1/2})$ of charge decay by roughly a factor of 10 at saturation, while particle composition, size (in the range of 0.5 to 2 mm), and polarity seem to have little effect. Asymmetries in the effect of UV radiation requires further study but can provide insight into the relative influence of radiation on particles in Earth's atmosphere. We present a simple ion recruitment model that can accurately predict decay times and curves below 60% to 70% RH. Together, our experiments and models suggest that electrostatic forces may significantly influence the residence times of particles with diameters of 10s of microns.

Results and Discussion

We measured the decay of surface charge on single particles lofted in the air using a TinyLev Acoustic Levitator (TAL) (32) housed in an environmental chamber. TAL is an opensource, 3D-printed instrument capable of manipulating micronto millimeter-sized particles in a noncontact manner (see the "Materials and methods" section; Figs. 1a and b). We modified TAL with the ability to measure charges on suspended particles with a resolution of 1 fC. Particles placed in the acoustic trap are initially charged by ionizing the air in the chamber with a high-voltage supply connected to a bundle of sharp carbon needles. Free ions and electrons rapidly adhere to all surfaces, including the isolated particle. Using a Gerdien tube condenser ion counter (AlphaLab, Inc), we estimate the ion density at the location of the levitating particle to be 10 to $15 \times 10^6/\mathrm{cm^3}$ during the charging period. The gas ion concentration decreases rapidly (on the order of seconds) after the ionizer is shut off, presumably due to collisions with grounded metallic surfaces in the environmental chamber. Charges on the isolated particle's surface, however, can only discharge across the gas–solid boundary.

TAL can hold a trapped particle static and move it along its primary axis (Fig. 1c; Fig. S1; Movie S1). We use this capability to perform noncontact charge measurements across extended time frames: a levitating particle is periodically (every 1 to 5 min) lowered into and raised out of an "acoustically transparent" Faraday cup (ATFC) to ascertain its surface charge. The ATFC is connected to charge-sensitive electrometer, which provides an overall sensitivity of 1 pC/V. A typical signal from the output stage of the electrometer during one measurement cycle is shown in Fig. 1d. The voltage difference ΔV between the initial baseline and the curve's minimum is proportional to the amount of charge on a particle's surface. Because the particle never touches a surface during the measurement process, charge loss occurs only through interactions with the gas. Additional details about the charge measurement system are included in the "Material and Methods" section and in Harper et al. (25, 35).

To approximate the diversity of particles suspended in Earth's atmosphere (from silicate dust to pollen to microplastics), we explored the decay of charge on particles of four compositions: EPS, silica aerogel, toasted amaranth (Perú), and volcanic pumice (ejecta from the 1932 Quizapú pumice, white Pumice from Popocatépetl, and pumice from the Kos Plateau Tuff). Exemplary grains are rendered photographically in Figs. 1e to g. All particles, despite a wide variation in shape, had spherical-equivalent diameters (SED) ranging between 1 to 2 mm. Although TAL can suspend particles with densities as high as 3,000 kg/m³, the device has difficulty moving such dense materials up and down as required for our charge measurement. Therefore, we did not conduct experiments with solid silica sand. Nonetheless, both the pumices and the aerogel have high silica contents (68 to 75% (36-38) and 100%, respectively), making them first-order analogs for silicate particles. Furthermore, as will be discussed later on, charge decay on isolated particles is largely independent of composition and morphology.

EPS particles were transferred to the experimental setup directly from their packaging container (nylon bag). Conversely, aerogel and pumice samples were crushed to the appropriate size and were then stored in a desiccator at < 10% RH until beginning an experiment. We toasted amaranth samples immediately prior to inserting them into TAL. Particles placed into TAL may be precharged as a result of contact electrification during handling. For instance, Fig. 2a shows that EPS particles taken from their plastic packaging bag generally carry positive charge [consistent with the relative positions of polystyrene and nylon on the triboelectric series (39)]. However, the ionization process is capable of erasing this initial bias (Fig. 2b; Fig. S2).

The environmental chamber which encloses TAL allowed us to explore the effect of atmospheric water (at 25°C) on particle charge decay. The decay dynamics for a subset of EPS particles at



Fig. 1. (a) A TinyLev acoustic levitator suspends a 2 mm polystyrene bead above the acoustically transparent Faraday cup. (b) Diagram of the charge measurement system inside the environmental chamber with a theoretical acoustic field. (c) Schlieren image of the acoustic field with a single levitated particle (33, 34). A second, virtual image of the particle is visible in the mirror (see Supplementary Material). (d) Output of the amplification stage as the particle was lowered into and raised out of the Faraday cup. ΔV that corresponds to the charge on the particle. (e) Expanded polystyrene bead (EPS). (f) Toasted amaranth grain. (g) Pumice from the 1932 eruption of the Quizapú volcano (Maule, Chile).



Fig. 2. (a) The probability density function of charge on EPS particles before ionization. Particles have some initial positive charge after being extracted from the container. (b) Distribution of charge on particles immediately after ionization with negative or positive bias. (c) Normalized voltage $\Delta V/\Delta V_0$ vs time for different RH (symbols). The low RH data are fit well by Eq. 2 as shown by the red line, which fits the 66% RH decay. The fit parameters are $Q_0 = 1$, $Q_c = 0.11$, and K = 1.24.

various RH are rendered in Fig. 2c. In dry environments, the functional form of the charge decay was linear at early times, giving way to nearly exponential behavior at long times (see black and red curves Fig. 2c). For RH < 30%, all particles experienced charge decay over a characteristic half-life, $t_{1/2}$, of 2 to 8 days. At higher RH, a more expedient decay rate followed a logistic form (green, pink, and purple curves). This timescale decreased to 1 day or less for RH approaching 100%. Yet, even the shortest decay times in our experiments greatly exceed those of seconds to hours reported in several previous works for broad ranges of RH (40, 41). We note that those efforts did not use isolated charged surfaces (i.e. tested materials were clamped, tethered, or otherwise secured extraneous surfaces). We suspect that shorter decay times there reflect conduction processes in addition to charge exchange between solid and gas. Indeed, Burgo et al. (42) found multi-day decay times for charged, polyethylene slabs in contact with aluminum, but short decay times at higher RH, possibly resulting from the growth of a conducting water layer on the surface.

Using the TAL system, particles only discharge by exchanging charge carriers with the atmosphere. Previous efforts have invoked ion recruitment processes to account for charge gain or loss on particle and aerosol surfaces (43, 44). The conventional prediction from such models is that particles reach electrostatic equilibrium in minutes to hours, not days. The dramatic departure from theory implied by the long charge decay times observed in our experiments may indicate that the regions near the charged object are depleted of free ions. Such depletion was recently suggested by Heinert et al. (44) to explain the multi-day decay of charge on a magnetically levitated conducting disk.

Particles in our experiments have radii of $a \approx 1 \,\mathrm{mm}$ and initial charges $|Q_0| \leq 1 \, \text{pC} = 6.25 \times 10^6$ elementary charges. These conditions correspond to approximately 0.5 charges for every square micron of particle surface area. While this charge density may seem dilute, the background ion density, as measured by the Gerdien tube condenser, is less than 1 ion/mm³ for both positive and negative ions. The dynamics of charge neutralization in our experiments can be understood by considering the ratio of the potential energy to thermal kinetic energy for a single airborne ion: $e\phi/k_{\rm B}T = eQ/4\pi\epsilon_0 r k_{\rm B}T$. In this framework, $k_{\rm B}$ is Boltzmann's constant, T is the temperature, ϵ_0 is the permittivity of free space, e is the elementary charge, and r is the separation between the ion and the charged particle. We make the simplifying assumption that oppositely-charged ions near the particle are always captured, whereas ions far from the particle can escape capture. The boundary between these regimes is estimated as $|e\phi/k_{\rm B}T| \approx 1$.

For the typical parameters described above when $|Q_0| = 1 \text{ pC}$, this boundary corresponds to a radius $r \approx 35 \text{ cm}$, which is larger than the size of our experimental chamber. Thus, in the early stages of charge decay, we can assume that every oppositely charged ion generated in the chamber is captured by the charged, millimetric particle. Indeed, the near linear decays seen in Fig. 2c for low RH suggest that discharge rates are likely determined by and equal to the rate of ion production in the air surrounding the particle. However, as the particle loses surface charge, the capture radius becomes significantly smaller than the size of the experimental chamber. For example, when $Q = 0.1Q_0$, $r \approx 3.5 \text{ cm}$, and the decay rate of charge will depend on Q. Moreover, free ions may be funneled away by the flow system that maintains the chamber at low RH or may be neutralized by ions of opposite polarity before being captured by the particle.

A very simple model for the charge decay that includes both of these regimes can be written as

$$\frac{dQ}{dt} = -\frac{K}{1+Q/Q_c}Q,\tag{1}$$

where K is a characteristic rate and Q_c is a characteristic charge representing a crossover between these regimes. A similar equation results from the capture of molecules by diffusion to a spherical particle covered in absorbing patches, as first discussed by Berg and Purcell in the context of chemoreception (45). This equation can be readily solved for Q(t):

$$Q(t) = Q_c W \left[\frac{Q_0 \exp\left(Q_0/Q_c - Kt\right)}{Q_c} \right],$$
(2)

where W represents the Lambert function and Q_0 is the initial charge on the particle. This three-parameter function shows excellent agreement with the data in dry conditions, as shown by the solid line in Fig. 2c. It is important to note that Eq. 2 does not depend on the size of the particle, which may explain similar decay times observed in experiments with much larger, centimeter-scale objects (42, 44).

Even though data for RH as high as 60% to 70% follows the same trend as for very dry conditions, real lofted particles are often found in water-rich environments near saturation. As Fig. 2c shows, the decay rate of charge increases significantly with higher RH. Generally, equilibrium ion concentrations increase with RH (46), but the rate at which a system recovers to its equilibrium concentration is not well-known. In our experiments, we assume the charged particle and surfaces of the TAL apparatus quickly deplete nearly all ions in the chamber once the ionizer is shut off. Thus, we expect that higher RH acts to return the ion con-



Fig. 3. $t_{1/2}$ vs RH for (a) different materials that were negatively charged and (b) EPS that was positively or negatively charged. We observed similar trends for $t_{1/2}$ with increasing RH for all materials used in (a). In (b), we observe a slightly smaller half life at lower RH for EPS particles, possibly due to differences in positive and negative ion concentrations.

centration to equilibrium more rapidly. Initially, the charge decay rate is small before increasing at later times as the ion concentration returns to equilibrium, resulting in a logistic-shaped decay curve, as seen in Fig. 2c. This particular decay, characterized by an increase in decay rate during the lifetime of the experiment, has been observed in other studies (42, 44). It is unclear if this behavior is solely related to RH, or other uncontrolled conditions.

Although data at higher RH cannot be fitted with Eq. 2, we can compare the half-life, $t_{1/2}$, of charge decay across all materials and RH. We note that there is a natural variability in decay time for individual experiments conducted at the same RH (see Fig. 2c). Nonetheless, the data in Fig. 3a shows no observable variation of t_{1/2} for the different particles used in our experiments. These particles vary in their shape, size, porosity, hydrophobicity, and spectral properties. Additionally, $t_{1/2}$ decreases with RH by roughly a factor of 10 at saturation for all particle types. This is consistent with our model since the capture probability of a given ion is only dependent on the atmospheric conditions and the total charge on the levitating particle. It is possible that in ion-rich environments, the charge decay of these particles would vary based on local surface properties or combinations of ion transport mechanisms such as diffusion, electrostatic drift, and convection. In our experiments, the acoustic levitator generates a small amount of heat, which nonetheless causes convection that is visible in the Schlieren imaging of the acoustic field (Movie S1). However, these environmental variations occur on much shorter timescales where we observe no large fluctuations in the rate of charge decay.

Atmospheric dust may carry a net positive or negative charge (24, 25, 47). Fig. 3b shows results for negatively and positively charged EPS particles, indicating little difference in $t_{1/2}$. If discharge is mostly determined by ion recruitment from the atmosphere, then the decay rate should represent the relative concentrations of positive and negative ions. Measurements suggest a roughly equal number of positive and negative ions in our exper-



Fig. 4. $\Delta V / \Delta V_0$ vs. time for particles irradiated by UV of different wavelengths, as indicated in the legend. Panel (a) represents negatively charged EPS and panel (b) represents positively charged EPS. The inset shows a zoom-in of positively charged particle data with the shortest decay times. The different symbols in the inset correspond to repeated experiments.

imental chamber, explaining the similarity between decay rates for positively and negatively charged particles.

Lofted particles in Earth's atmosphere also experience varying amounts of UV radiation, ranging from UVC at the highest altitudes, to UVA near the Earth's surface. To simulate this exposure, in our experiments we used two different types of UV light bulbs. Naively, one may expect that UV radiation always accelerates charge decay as high-energy photons neutralize surface charge. Previous studies have shown that the efficiency of UV light on discharging can vary dramatically in a narrow range of wavelength (48). Strikingly, we observe that UV radiation can even extend the life of charge on some particles. Fig. 4 shows data for EPS particles with negative (4a) and positive (4b) initial charge in dry conditions (low RH). With no UV radiation, particles exhibit a multi-day decay timescale $(t_{1/2})$, as expected. For UVA radiation (365 nm, 3.40 eV), $t_{1/2}$ decreased slightly to 2 to 3 days for negatively charged particles, yet $t_{1/2}$ increased significantly for positively charged particles.

This asymmetry can be explained by noting that UV radiation can potentially ionize organic atmospheric impurities that could deplete the local ion concentration, or even by direct photoelectric charging of the particle. The charged surface states can be highly variable given the complex particle materials, and the potential presence of water. Our experiments cannot disentangle these effects at the moment, however, the asymmetry is highly relevant for most airborne particles since UVA radiation can penetrate to Earth's surface. For UVC radiation (254 nm, 4.88 eV), which is absorbed in Earth's upper atmosphere, particles of either charge polarity experienced a decrease in $t_{1/2}$, yet positively charged particles decayed in less than 1 hour. This drastic decrease in the decay rate is almost certainly due to photoelectric electrons emitted from the surface of the copper mesh comprising the Faraday cup. The work function of copper is 4.7 eV.

How does the multi-day decay of charge affect transport in the atmosphere? The dynamics of a particle with a spherical equivalent diameter (SED) of 2*a* settling out of the atmosphere are governed by gravitational, F_g , drag, F_d , and electrostatic forces, F_e . Here, we assume the simplest 1D model of this process (see Supplementary Material for full treatment):

$$m\dot{v} = F_g + F_d + F_e, \tag{3}$$



Fig. 5. Model results for silicate particle sedimentation times assuming no turbulent mixing from an altitude of 5 km. (a) Without electrostatic forces, t_{set} is determined by gravitational and drag forces, ranging from years to hours for particle diameters 1 to 100 μ m. (b) Percent change in settling time for a positively charged particle in an external electric field with a surface charge density $\sigma = 10^{-5}$ C/m² that decays exponentially with t_{1/2} = 4 days. The symbols represent different ambient electric fields (+, upward pointing; – downward pointing).

where *m* is the particle mass and \dot{v} is the acceleration in the direction of gravity. We can quantify the impact of charge decay on atmospheric residence time by computing the difference sedimentation time t_{set} between a charged particle and an otherwise identical neutral particle. For this analysis, we consider particles with SEDs in the range of 1 to 100 μ m and densities between 1,000 and 2,900 kg/m³. Earth's fair weather electric field is on the order of 0.1 kV/m and points toward the surface. However, ~10 kV/m is typical in dust storms (31), and \sim 100 kV/m has been measured during foul weather (49). As such, we employ a conservative electric field range spanning ±5 kV/m. Furthermore, we assume that particles are initially charged to the theoretical maximum limit of $\sigma \sim$ 10^{-5} C/m² and then decay exponentially with a half-life of 4 days. Lastly, we determined settling times for particles falling from an altitude of 5 km, an elevation at which dust has been detected in the Saharan Air Layer (50). The model implements a simple atmospheric profile to account for changing pressure as a particle descends (Fig. S3). We note that our model excludes the effects of turbulence and advection.

Settling times from an altitude of 5 km can range from a few hours for the largest particles to $>10^2$ days for the smallest grains (see Fig. 5a and Fig. Supplementary Material). Because small, neutral particles take months or even years to settle, electrostatic effects with half-lives of a few days should not affect tset. Likewise, electrostatic forces play a minimal role for large particles since these are dwarfed by inertial forces. Conversely, our numerical experiments show that when $t_{set} \approx t_{1/2}$, electrostatic effects can substantially modify particle residence times. Indeed, we find that charge most strongly influences the dynamics of 2,900 kg/m³ particles (mineral dust) with SEDs of 5 to $20 \,\mu$ m by increasing t_{set} by 80% (Fig. 5b). For lighter 1,000 kg/m³ particles (e.g. microplastics), the effects of electrostatic forces are shifted toward particles with SEDs of 10 to $30\,\mu\text{m}$ and can modify the settling times by up to 175% (Fig. S4). These results are consistent with both experimental (21) and inferences from field observations (20).

Beyond more complex fluid dynamics, we note that the above model neglects additional electrostatic processes that may retard, reverse, or accelerate charge loss. As aforementioned, charged dust plumes have been detected at large distances from their sources, suggesting that in situ charging mechanisms keep lofted particles electrified well beyond the timescales measured here (28, 50). Additionally, we suspect that other environmental factors (such as pressure and temperature) may influence the retention of charge, and these should be explored in future work. However, our first-order analysis highlights the effects electrostatic forces may have on the transport of particles with $2a \sim 10 \,\mu$ m. Whether these forces help retain particles in the atmosphere (or more easily remove them) requires a better understanding of the charge distribution in dust layers and electric fields at elevation.

Conclusions

Using acoustic levitation and noncontact charge measurement, we find that isolated particles of any material can retain charge for weeks. In general, RH decreased the half-life of charge $(t_{1/2})$, but the effect was less pronounced than reported in previous works (40, 42). Additionally, $t_{1/2}$ was insensitive to the charge polarity of the particles for a broad range of RH. However, exposure to UV light either accelerated or arrested charge decay, depending on charge polarity and UV wavelength. Both RH and irradiation control the local concentration of ions in the air surrounding a charge particle, and ultimately determine the longevity of charge on a particle's surface. Using our experimental data as input to a simple 1D model, we find that electrostatic forces significantly modulate the residence times of airborne coarse mode dust with diameters 5 to 30 μ m. We provide the first measurements of charge decay on mm-scale lofted particles, and demonstrate that electrostatic forces and charge decay must be considered in dust transport models.

Materials and Methods Environment setup

The RH of the chamber was varied by combining boiling water and dry building air. RH values were maintained within 10% variance over the entire experiment, with the majority of data points within 5% RH variance. For data presented in Fig. 3, the environment was set before charging the particle. For UV experiments, the UV bulbs were suspended at the inside wall of the chamber, horizontally aligned with the particle. The chamber was left partially open to laboratory air, and the initial ΔV_0 was measured before UV irradiation. The 254 nm bulb produced irradiance at the particle's position of $1.06 \times 10^{-1} \text{ W/m}^2$. The 365 nm bulb had irradiance of $7.07 \times 10^{-3} \text{ W/m}^2$ at the particle's position.

Acoustic levitation

The acoustic levitator follows the open-source "TinyLev" design presented in Marzo et al. (32). TinyLev consists of two semispherical arrays of 40 kHz transducers separated by a distance of 11.5 cm. Generally, both hemispheres are driven in phase, creating a standing wave that holds a particle static. However, by slightly changing the phase of one hemisphere, a particle in the acoustic trap may be moved up or down. We utilize this technique to lower and raise the particle in and out of the Faraday cage in a noncontact manner.

Charge measurement

Levitated particles are charged for 10s with a corona ionizer attached to a Bertan Series 225 high-voltage power supply at \pm 8 to 10kV to eliminate any initial charge bias (Figs. 2a and S2). After charging, a 5 min wait period allows charge to dissipate from surfaces in the chamber. We measured the particle charge by acoustically lowering it by 2 λ into the ATFC, then raising it back to its starting position. This was done slow enough to maintain the integrity of the acoustic trap. The ATFC is connected to an electrometer consisting of a charge-sensitive preamplifier and an inverting amplifier stage with a gain of 100. Thus, the electrometer generates voltage pulses that are proportional to the charge on a particle entering the ATFC according to

$$\Delta V = -100 \times \frac{Q}{C} e^{-t/RC}.$$
 (4)

Here, the time constant of the charge amplifier's feedback loop was RC = 5 s, where C = 1 nF and $R = 5 G\Omega$.

For most experiments, we performed charge measurements every 1 to 5 min. The rate was increased to 2 measurements/minute for experiments in which particle charge decayed rapidly (high humidity or UV). This process was repeated indefinitely until the end of the experiment.

Acknowledgments

The authors thank the anonymous reviewers for their valuable suggestions. We thank Joe Dufek for insightful conversations.

Supplementary Material

Supplementary material is available at PNAS Nexus online.

Funding

This material is based upon work supported by the National Science Foundation (NSF: # 2010524).

Authors' Contributions

J.M.H. and D.H. contributed equally to this work and should be considered first coauthors. J.M.H. designed, implemented, and executed experiments, as well as performed data analysis. D.H. executed experiments and performed data analysis. T.H. aided in construction of experimental apparatus, conducted experiments, and performed data analysis. J.M.I.I.I. aided in the construction of experiments and conducted experiments. D.M. performed schlieren imaging experiments. J.C.B. performed experiments, data analysis, and modeling.

Preprints

A preprint of this article is published at https://arxiv.org/abs/2206 .06848.

Data Availability

All data are included in the manuscript and/or Supplementary Material.

References

- Kinne S, et al. 2006. An aerocom initial assessment—optical properties in aerosol component modules of global models. Atmos Chem Phys. 6(7):1815–1834.
- Kok JP, et al. 2017. Smaller desert dust cooling effect estimated from analysis of dust size and abundance. Nat Geosci. 10:274– 278.
- Kanji ZA, et al. 2017. Overview of ice nucleating particles. Meteorol Monogr. 58(1):1.1–1.33.
- Young CL, Sokolik IN, Flanner MG, Dufek J. 2014. Surface radiative impacts of ash deposits from the 2009 eruption of redoubt volcano. J Geophys Res Atmos. 119(11):11387–11397.
- Kaspari S, Skiles M, Delaney I, Dixon D, Painter TH. 2015. Accelerated glacier melt on Snow Dome, Mount Olympus, Washington, USA, due to deposition of black carbon and mineral dust from wildfire. J Geophys Res Atmos. 120(7):2793–2807.
- Méndez Harper J, Steffes P, Dufek J, Akins A. 2019. The effect of electrostatic charge on the propagation of GPS (L-band) signals through volcanic plumes. J Geophys Res Atmos. 124(4):2260– 2275.
- Tegen I, Lacis AA. 1996. Modeling of particle size distribution and its influence on the radiative properties of mineral dust aerosol. J Geophys Res Atmos. 101(D14):19237–19244.
- Idso SB, Brazel AJ. 1977. Planetary radiation balance as a function of atmospheric dust: climatological consequences. Science. 198(4318):731–733.
- Chaibou AAS, Ma X, Sha T. 2020. Dust radiative forcing and its impact on surface energy budget over West Africa. Sci Rep.10(12236). https://www.nature.com/articles/s41598-020 -69223-4.
- Ryder CL, et al. 2019. Coarse and giant particles are ubiquitous in Saharan dust export regions and are radiatively significant over the Sahara. Atmos Chem Phys. 19(24):15353–15376.
- Yu H, et al. 2015. The fertilizing role of African dust in the Amazon rainforest: a first multiyear assessment based on data from Cloud-Aerosol Lidar and Infrared Pathfinder Satellite Observations. Geophys Res Lett. 42(6):1984–1991.
- 12. Usher CR, Michel AE, Grassian VH. 2003. Reactions on mineral dust. Chem Rev. 103(12):4883–4940.
- Thomas P, Gierasch PJ. 1985. Dust devils on mars. Science. 230(4722):175–177.
- Méndez Harper J, Helling C, Dufek J. 2018. Triboelectrification of KCI and ZnS particles in approximated exoplanet environments. Astrophys J. 867(2):123.
- 15. Rodriguez S, et al. 2018. Observational evidence for active dust storms on Titan at equinox. Nat Geosci. 11(10):727–732.
- McDonald GD, et al. 2022. Aeolian sediment transport on io from lava-frost interactions. Nat Commun. 13(2076). https://www.na ture.com/articles/s41467-022-29682-x.
- 17. Weinzierl B, et al. 2017. The saharan aerosol long-range transport and aerosol-cloud-interaction experiment: Overview and selected highlights. Bull Amer Meteor Soc. 98(7): 1427-1451.
- Adebiyi AA, Kok JF. 2020. Climate models miss most of the coarse dust in the atmosphere. Sci Adv.6(15):9507.
- Denjean C, et al. 2016. Size distribution and optical properties of mineral dust aerosols transported in the western Mediterranean. Atmos Chem Phys. 16(2):1081–1104.
- Maring H, Savoie DL, Izaguirre MA, Custals L, Reid JS. 2003. Mineral dust aerosol size distribution change during atmospheric transport. J Geophys Res Atmos.108(D19). https://agupubs.onlinelibrary.wiley.com/doi/epdf/10.1029/2002JD002536.

- Toth III JR, et al. 2020. Electrostatic forces alter particle size distributions in atmospheric dust. Atmos Chem Phys. 20(5):3181– 3190.
- 22. Kok JF, Renno NO. 2008. Electrostatics in wind-blown sand. Phys Rev Lett. 100(1):014501.
- Cimarelli C, Behnke S, Genareau K, Méndez Harper J, Van Eaton A. 2022. Volcanic electrification: recent advances and future perspectives. Bull Volc. 84(8):1–10.
- James MR, Lane S, Gilbert JS. 2000. Volcanic plume electrification: experimental investigation of a fracture-charging mechanism. J Geophys Res Solid Earth. 105(B7):16641–16649.
- Méndez-Harper J, Cimarelli C, Cigala V, Kueppers U, Dufek J. 2021. Charge injection into the atmosphere by explosive volcanic eruptions through triboelectrification and fragmentation charging. Earth Planet Sci Lett. 574(14):117162.
- Aplin K, Nicoll K, Houghton I. 2015. In: Electrical charging of volcanic ash from Eyjafjallaj. EGU General Assembly Conference Abstracts. p. 11712. https://ui.adsabs.harvard.edu/abs/20 15EGUGA..1711712A.
- 27. Okuzumi S. 2009. Electric charging of dust aggregates and its effect on dust coagulation in protoplanetary disks. Astrophys J.698(2):1122.
- Harrison RG, Nicoll KA, Ulanowski Z, Mather TA. 2010. Selfcharging of the Eyjafjallajökull volcanic ash plume. Environ Res Lett. 5(2):024004.
- Harrison RG, Nicoll KA, Marlton GJ, Ryder CL, Bennett AJ. 2018. Saharan dust plume charging observed over the UK. Environ Res Lett. 13(5):054018.
- Méndez Harper JS, et al. 2017. Electrification of sand on Titan and its influence on sediment transport. Nat Geosci. 10(4):260–265.
- Zhang H, Zhou Y. 2020. Reconstructing the electrical structure of dust storms from locally observed electric field data. Nat Commun.11(5072):1–12.
- Marzo A, Barnes A, Drinkwater BW. 2017. TinyLev: A multiemitter single-axis acoustic levitator. Rev Sci Instrum.88(8). 085105.
- Settles GS, Covert EE. 2002. schlieren and shadowgraph techniques: visualizing phenomena in transport media. Appl Mech Rev. 55(4):B76–B77.
- Crockett A, Rueckner W. 2018. Visualizing sound waves with schlieren optics. Am J Phys. 86(11):870.
- Méndez Harper J, Courtland L, Dufek J, McAdams J. 2020. Microphysical effects of water content and temperature on the triboelectrification of volcanic ash on long time scales. J Geophys Res Atmos. 125(14):e2019JD031498.
- Allen SR. 2001. Reconstruction of a major caldera-forming eruption from pyroclastic deposit characteristics: Kos Plateau Tuff, eastern Aegean Sea. J Volc Geoth Res. 105(1-2):141–162.
- Siebe C, et al. 2017. The ~23,500 y 14c bp white Pumice Plinian eruption and associated debris avalanche and Tochimilco lava flow of Popocatépetl volcano, México. J Volc Geoth Res. 333:66– 95.
- Ruprecht P, Bergantz GW, Cooper KM, Hildreth W. 2012. The crustal magma storage system of Volcán Quizapu, Chile, and the effects of magma mixing on magma diversity. J Petrol.53(4):801– 840.
- Zou H, et al. 2019. Quantifying the triboelectric series. Nat Commun.10(1427):1–9.
- Tada Y, Murata Y. 1995. Direct charge leakage through humid air. Jpn J Appl Phys.34:1926–1927.
- Kumara S, Serdyuk YV, Gubanski SM. 2011. Surface charge decay on polymeric materials under different neutralization modes in air. IEEE Trans Dielectr Electr Insul.18(5):1779–1788.

- Augusto de Lima Burgo T, Rezende AC, Bertazzo S, Galembeck A, Galembeck F. 2011. Electric potential decay on polyethylene: role of atmospheric water on electric charge build-up and dissipation. J Electrostat. 69(4):401–409.
- Gunn R. 1954. Diffusion charging of atmospheric droplets by ions, and the resulting combination coefficients. J Atmos Sci. 11(5):339–347.
- 44. Heinert C, Sankaran RM, Lacks DJ. 2022. Decay of electrostatic charge on surfaces due solely to gas phase interactions. J Electrostat. 115:103663.
- 45. Berg HC, Purcell EM. 1977. Physics of chemoreception. Biophys J. 20(2):193–219.

- Carlon HR. 1981. Equilibrium ion content of water vapor in air. J Appl Phys. 52(4):2638.
- 47. Jungmann F, Kruss M, Teiser J, Wurm G. 2022. Aggregation of sub-mm particles in strong electric fields under microgravity conditions. Icarus. 373:114766.
- Ugolini D, Girard M, Harry GM, Mitrofanov VP. 2008. Discharging fused silica test masses with ultraviolet light. Phys Lett A. 372(36):5741–5744.
- Dwyer JR, Uman MA. 2014. The physics of lightning. Phys Rep. 534(4):147–241.
- 50. Nicoll KA, Harrison RG, Ulanowski Z. 2010. Observations of saharan dust layer electrification. Environ Res Lett. 6(1):014001.